

# Near-surface bedrock profiling using urban ambient noise: an autocorrelation approach

## INTRODUCTION

#### **Challenges :**

- Near-surface bedrock may vary from a few meters to a few tens of meters underground.
- Civil engineers' common practice for bedrock profiling is drilling boreholes.
- Because of strict regulations in an urban environment, non-invasive geophysical surveys are demanded to complement boreholes drilling.
- Civil engineering projects require very high resolution of a few meters both vertically and horizontally.

#### Background:

- Interferometry theory turns continuous ambient noise to empirical Green's functions by cross-correlating any two receivers' recordings.
- To the extreme, crosscorrelation becomes autocorrelation.
- Claerbout [1985] pointed out that "by autocorrelating the data of hours and days duration we convert the chaos of continuing microseismic noise to something that might be the impulse response of the earth".
- Some recent studies showed the feasibility to retrieve P-wave reflection and image deep earth structure by autocorrelating microseismic noise [Ito and Shiomi, 2012; Saygin et al., 2017 and Taylor et al., 2016].

## **TUNING PROCESS**

### Autocorrelation of ambient noise

- Based on the idea of Claerbout [1985], autocorrelation of ambient noise  $U(x_i, t)$  can approximate a zero-offset seismic trace.
- It follows the convolution model and can be expressed as:

 $U(x_i,t) * U(x_i,t) \approx w(x_i,t) \otimes r(x_i,t) \otimes e(x_i,t),$ 

where  $w(x_i, t)$ : source signature,  $r(x_i, t)$ : receiver signature, and  $e(x_i, t)$ : reflectivity.

### Assumptions

- Source signature is identical within a certain area:  $w(x_i, t) \rightarrow w(t)$ .
- Identical and negligible receivers' effect:  $r(x_i, t) \rightarrow r(t)$ .

### Required prior knowledge

• Bedrock depth from a reference borehole is required to estimate a two-layer reflection coefficient  $\hat{e}(x_{ref}, \omega)$ .

#### Coupled signature of source and receiver

$$\widehat{w}(\omega) \ \hat{r}(\omega) \approx \frac{U(x_{ref},\omega) * U(x_{ref},\omega)}{\hat{e}(x_{ref},\omega)}$$

Approximated reflectivity

$$e(x_i, t) = \mathcal{F}^{-1}\hat{e}(x_i, \omega) = \mathcal{F}^{-1} \frac{U(x_i, \omega) * U(x_i, \omega)}{\widehat{w}(\omega) \hat{r}(\omega)}$$

#### Time to depth conversion

• Require another reference borehole to estimate average velocity of soil layer.  $2(d_1 - d_2)$ 

$$V = \frac{2(u_1 \ u_2)}{t_1 - t_2}$$
,

where  $d_1$  and  $d_2$  are bedrock depths based on the two reference boreholes.

• Bedrock depths at other traces can be estimated:  $d_i = d_1 + 0.5\overline{V} (t_i - t_1)$ 





<sup>1</sup>Department of Civil & Environmental Engineering, National University of Singapore <sup>2</sup>Chinese Academy of Sciences, Institute of Tibetan Plateau Research Email: zhangyunhuo@u.nus.edu, ceeliyy@nus.edu.sg, zhangheng415@itpcas.ac.cn, ceekt@nus.edu.sg.

## Yunhuo Zhang<sup>1</sup>, Yunyue Elita Li<sup>1</sup>, Heng Zhang<sup>2</sup> & Taeseo Ku<sup>1</sup>